



The combined processing of geomagnetic intensity vector projections and absolute magnitude measurements

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Abstract. A method for the combined processing measurements of the projections and the absolute magnitudes of geomagnetic field intensity vectors, based on mathematical technology of local approximation models, was proposed. The approach realized in this paper, based on the proposed method, provides the increase of accuracy of the measurements of the projections of the geomagnetic intensity vector. An algorithm for the two-stage combined processing of the measurements of

15 projections and absolute magnitudes of geomagnetic field intensity vectors was developed. The operation of the combined processing algorithm was tested on models and observatory measurements. The estimates of the combined processing algorithm errors were obtained using statistical modelling. The reduction of the root mean square error values was achieved for the estimates of the projections of geomagnetic field intensity vectors.

1 Introduction

- 20 In this manuscript, a method and an algorithm for combined processing measurements of the components (projections and the absolute magnitudes) of geomagnetic field intensity vectors is proposed. The approach realized in the paper, based on the proposed method, provides the increase of accuracy of the measurements of the projections of the geomagnetic intensity vector. The considered measurements are carried out by INTERMAGNET observatories equipped with systems of vector and scalar magnetometers; the definitive type data are used, containing systematic errors equal to zero (Geomagnetic, 2011;
- 25 INTERMAGNET, 2018). The measurement errors of vector and scalar magnetometers are represented by random normally distributed errors with zero expectation and a predetermined variance. As usual, the measurement errors of the projections of geomagnetic field intensity vectors are significantly larger than the ones of the absolute vector magnitudes performed by the mentioned measurement devices. The formulation of the problem of reducing the noise root-mean-square (RMS) error values in the geomagnetic field intensity projection measurements due to combined processing of the values of all its
- 30 components.

In the following research:





1. a method for combined processing of the measurements of projections and absolute magnitudes of geomagnetic field intensity vectors is formulated, based on formation of the sequences of piecewise-constant local models followed by their weighted averaging;

a two-stage algorithm for combined processing of the measurements of projections and absolute magnitudes of
 geomagnetic field intensity vectors is developed;

3. the testing of the algorithm on model and observatory data is performed;

4. the estimates of the algorithm errors, calculated using statistical modelling, are presented; the reduction of the RMS noise errors for the estimates of the geomagnetic field vector projections is proved.

- The materials of this research paper are intended for specialists (magnetologists) engaged in tasks of digital processing of geomagnetic field measurements. The need to reduce noise errors in estimates of the projections of the geomagnetic field intensity vectors measured by vector magnetometers arises in a number of technical and scientific applications. For instance, technogenic disturbances can affect the geomagnetic observatory hardware, affecting vector and scalar magnetometers differently: as a rule, the noise errors from vector magnetometers occurring due to such interference are greater than the noise errors from scalar magnetometers. The decrease of noise errors from vector magnetometers is necessary, for example,
- 15 for calculation of the gradients of the projections of the geomagnetic field intensity vectors in the navigation problems. Nowadays the reduction of errors for vector magnetometers (with certain assumptions) is achieved using optimization of the calibration from scalar magnetometers (Merrayo et al., 2000; Olsen et al., 2003) or refinement of calibration characteristics (Soborov et al., 2008) based on special mathematical processing. In the measurement systems considered in this paper, in fact, parallel measurements are performed; a possible algorithm providing the decrease of errors for such measurements can
- 20 be formed based on Kalman filters (Shakhtarin, 2008). However, due to the peculiarities of this problem, the construction of the resulting nonlinear filters is associated with certain problems due to the inaccuracies of linearization and the accepted hypothesis concerning the type of initial intensity vector function. In the research (Soloviev et al., 2018), joint processing of vector and scalar magnetometer measurements is aimed at improving the calibration accuracy of the so-called baseline, which only indirectly provides the considered reduction in errors. The combined processing of measurements of projections
- 25 and absolute magnitudes of geomagnetic field intensity vectors proposed in this paper is significantly free from the mentioned problems.

2 A method for combined processing of measurements of projections and absolute magnitudes of geomagnetic field intensity vectors

Let $H_1(Ti)$, $H_2(Ti)$, $H_3(Ti)$ and $H_0(Ti) = \sqrt{H_1^2(Ti) + H_2^2(Ti) + H_3^2(Ti)}$ be the initial functions for the projections and 30 absolute magnitudes of the geomagnetic field intensity vectors; we assume that $Y_1(Ti)$, $Y_2(Ti)$, $Y_3(Ti)$, $Y_0(Ti)$ are their values registered by vector and scalar magnetometers, $i = 0, 1, ..., N_f - 1$, the sampling interval T = 1 second; 1-second measurements from INTERMAGNET observatories are analyzed in this study. We represent the noise errors of the





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measurement values $W_n(Ti)$, n = 0, ...,3, in the form of uncorrelated, normally distributed random values with zero mathematical expectations and variances. Such representation of errors is, to a large extent, valid for cases of large technogenic noises that can occur when geomagnetic measurements are carried out. We consider, with some assumptions, that the spectrum of random components for the functions of the geomagnetic field is concentrated almost entirely in the low-frequency domain; the spectrum of random measurement errors is concentrated in the high-frequency domain.

We assume that the measurements, the initial functions and the errors are related by linear additive dependences:

 $Y_n(Ti) = H_n(Ti) + W_n(Ti), n = 0, ..., 3.$

Using the specified observation values $Y_n(Ti)$ we demand the determination of the estimates $Y_1^{\circ}(Ti)$, $Y_2^{\circ}(Ti)$, $Y_3^{\circ}(Ti)$, i = 0, 1, ..., $N_f - 1$, that would be close to the initial functions of the intensity vector projections. We perform the combined processing for the projections and magnitudes of the geomagnetic field intensity vectors in two stages.

10 At the first stage, on the main interval with the points $i = 0, 1, ..., N_f - 1$, we introduce the *N*-point sliding local intervals with limiting points N_{1_i}, N_{2_i} , and with the sliding step N_d and the quantity of sliding intervals m_0 .

$$N_{1_{j}} = N_{d}(j-1), \ N_{2_{j}} = N_{1_{j}} + N - 1, j = 1, \dots, m_{0}.$$
(1)

To simplify the considerations, we require the relations of multiplicity $mN = N_f$ and $N_d m_d = N$, here m, m_d are integers, then $m_0 = m_d(m-1) + 1$.

- For a sliding interval with a number *j* we define the model functions of a form $Y_{M_1}(c_{1j}, Ti)$, $Y_{M_2}(c_{2j}, Ti)$, $Y_{M_3}(c_{3j}, Ti)$, here c_{nj} are the vectors of model parameters, n = 1, 2, 3. These model functions can be, in particular, polynomial, piecewise constant, piecewise linear etc. The size of local intervals determines the approximation errors. At small *N* there will be large fluctuation errors, and at large *N* there will be large systematic approximation errors.
- Based on the above-defined measured values, models and the maximum likelihood method (Kramer 1975), we define the 20 local functional $S(c_j, Y_j)$ which determines the measure of closeness for local measurements and models, similar to (Getmanov., 2013) as the sum of the four functionals:

$$S_{0}\left(c_{1j}, c_{2j}, c_{3j}, Y_{0j}\right) = \sum_{i=N_{1j}}^{N_{2j}} \left(Y_{0j}(Ti) - \sqrt{\sum_{r=1}^{3} Y_{M_{r}}^{2}\left(c_{rj}, Ti\right)}\right)^{2},$$

$$S_{n}\left(c_{nj}, Y_{nj}\right) = \sum_{i=N_{1j}}^{N_{2j}} \left(Y_{nj}(Ti) - Y_{M_{n}}\left(c_{nj}, Ti\right)\right)^{2}, n = 1, 2, 3,$$

$$S(c_{j}, Y_{j}) = S_{0}\left(c_{1j}, c_{2j}, c_{3j}, Y_{0j}\right) + \sum_{n=1}^{3} S_{n}\left(c_{nj}, Y_{nj}\right).$$
(2)
Here $c_{j}^{T} = (c_{1j}^{T}, c_{2j}^{T}, c_{3j}^{T}), Y_{j}^{T} = (Y_{1j}^{T}, Y_{2j}^{T}, Y_{3j}^{T})$ are the parameter and value vectors related to the j^{th} local interval. Taking

25 into account the assumption of errors in measurements, we carry out the identification of the optimal estimates of the model parameters c_i° using the solutions of the sequence of optimization problems for local functionals

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$$c_{j}^{\circ} = \arg\left\{\min_{c_{j}} S(c_{j}, Y_{j})\right\}, j = 1, ..., m_{0}, c^{\circ^{T}} = (c_{1}^{\circ^{T}}, c_{2}^{\circ^{T}}, c_{3}^{\circ^{T}}).$$
(3)

We perform the construction of sliding local models in an obvious way by assuming $n = 1, 2, 3, j = 1, ..., m_0$.

$$Y_{n_j}^{\circ}(Ti) = Y_{M_n}\left(c_{n_j}^{\circ}, Ti\right), N_{1_j} \le i \le N_{2_j}, Y_{n_j}^{\circ}(Ti) = 0, 0 \le i < N_{1_j}, N_{2_j} < i \le N_f - 1$$
(4)

5 At the second stage we introduce the unit functions $E_i(Ti)$, equal to zero outside of local sliding intervals, add them together to get E(Ti) and calculate the sequence of weighting coefficients R(Ti)

$$E_{j}(Ti) = 1, N_{1_{j}} \le i \le N_{2_{j}}, \quad E_{j}(Ti) = 0, 0 \le i < N_{1_{j}}, \quad N_{2_{j}} < i \le N_{f} - 1$$

$$E(Ti) = \sum_{j=1}^{m_{0}} E_{j}(Ti), R(Ti) = 1/E(Ti), i = 0, \quad 1, \quad \dots, N_{f} - 1.$$
(5)

Let us perform the weighting averaging using Eq. (5) for the sum of the sliding local model sequence (Getmanov et al., 2015)

$$Y_n^{\circ}(Ti) = R(Ti) \sum_{j=1}^{m_0} Y_{n_j}^{\circ}(Ti).$$
(6)

We calculate the estimates $Y_n^{\circ}(Ti)$, $i = 0, 1, ..., N_f - 1, n = 1, 2, 3$, based on linear operations in Eq. (6) corresponding to the second stage.

The method of combined two-step processing of the values of the projections and the absolute magnitudes of geomagnetic

field intensity vectors consists in the sequential execution of the first and second stages in accordance with the Eq. (2), (3), 15 (4) and (5), (6).

3 An algorithm for two-stage processing of measurements of projections and absolute magnitudes of geomagnetic field intensity vectors for piecewise-constant models

Let us build the local intervals using Eq. (1) and define the local models on them as piecewise constant functions $Y_{M_n}(c_{n_i})$

 $Ti) = c_{n_i}, j = 1, ..., m_0, N_{1_i} \le i \le N_{2_i}$. Obviously, in this case, at local intervals with duration NT the initial functions for 20 geomagnetic field vector projections must be approximately constant; the local interval can be expanded if we treat piecewise constant functions as local models. Let us formulate the equation for local functionals:

$$S(c_j, Y_j) = \sum_{n=1}^{3} \left(\sum_{i=N_{1j}}^{N_{2j}} \left(Y_{n_j}(T_i) - c_{n_j} \right)^2 + \sum_{i=N_{1j}}^{N_{2j}} \left(Y_{0j}(T_i) - \sqrt{c_{1j}^2 + c_{2j}^2 + c_{3j}^2} \right)^2 \right)$$
(7)

We differentiate Eq. (7) with respect to c_{n_i} , equate the derivatives to zero and get the necessary conditions for an extremum 25 in the form of a system of three nonlinear algebraic equations as a result:

$$\frac{\partial s(c_j,Y_j)}{\partial c_{n_j}} = \sum_{i=N_{1_j}}^{N_{2_j}} Y_{n_j}(Ti) - 2Nc_{n_j} + \frac{c_{n_j} \sum_{i=N_{1_j}}^{N_{2_j}} Y_{0_j}(Ti)}{\sqrt{c_{1_j}^2 + c_{2_j}^2 + c_{3_j}^2}} = 0, n = 1, 2, 3.$$





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In this case, an exact analytical solution of this system is possible. Omitting the calculations, we obtain expressions for local estimates, $j = 1, ..., m_0$

$$c_{n_{j}}^{\circ} = \frac{1}{2N} \left(1 + \frac{\sum_{r=N_{1_{j}}}^{N_{2_{j}}} Y_{0_{j}}(Tr)}{\sqrt{\left(\sum_{r=N_{1_{j}}}^{N_{2_{j}}} Y_{1_{j}}(Tr)\right)^{2} + \left(\sum_{r=N_{1_{j}}}^{N_{2_{j}}} Y_{2_{j}}(Tr)\right)^{2} + \left(\sum_{r=N_{1_{j}}}^{N_{2_{j}}} Y_{3_{j}}(Tr)\right)^{2}}} \right) \sum_{i=N_{1_{j}}}^{N_{2_{j}}} Y_{n_{j}}(Ti), n = 1, 2, 3.$$

$$(8)$$

and get the piecewise constant functions for local estimates $Y_n^{\circ}(Ti) = c_{n_j}^{\circ}$, $N_{1_j} \le i \le N_{2_j}$, $Y_n^{\circ}(Ti) = 0$, $0 \le i < N_{1_j}$, $N_{2_j} < i \le N_f - 1$, n = 1, 2, 3, using Eq.(8) according to Eq.(4); let us present them as the realization of the first stage.

Weighted averaging of sequences of piecewise constant local estimates and the calculations of the estimate functions $Y_n^{\circ}(Ti)$ for the second stage are performed using Eq. (5), (6).

4 Testing of the algorithm for combined processing on model and observatory measurements

4.1 Testing on model data

10 The developed algorithm for combined processing was tested on model measurements. Initial model functions for the projections of the geomagnetic intensity vector $H_{MG_n}(Ti)$ were presented as quadratic functions; the model function for the absolute vector magnitude $H_{MG_0}(Ti)$ was calculated based on $H_{MG_n}(Ti)$, n = 1, 2, 3.

$$H_{MG_n}(Ti) = a_{0,n} + a_{1,n}Ti + a_{2,n}(Ti)^2, H_{MG_0}(Ti) = \sqrt{H_{MG_1}^2(Ti) + H_{MG_2}^2(Ti) + H_{MG_3}^2(Ti)},$$

 $i = 0, 1, ..., N_f - 1$. The noise error functions $W_n(Ti), n = 0, 1, 2, 3$, were modelled using normally distributed values with zero mathematical expectation and the variances $\sigma_1^2 = \sigma_2^2 = \sigma_3^2 = \sigma^2$ and σ_0^2 . The model values of the projections the absolute magnitude of the intensity vector were represented as sums:

$$H_{M_n}(Ti) = H_{MG_n}(Ti) + W_n(Ti), i = 0, 1, ..., N_f - 1, n = 0, 1, 2, 3$$
(9)

The following parameter values for the model functions were assigned: $a_{0,1} = 30000$, $a_{1,1} = 0.4206$, $a_{2,1} = -0.0040$, $a_{0,2} = 4300.4$, $a_{1,2} = a_{1,2} = -0.4252$, $a_{2,2} = 0.0043$, $a_{0,3} = 4089.5$, $a_{1,3} = 0.5013$, $a_{2,3} = -0.0056$, T = 1 sec. Model measurement values $H_{M_n}(Ti)$ were on the input of the Eq. (8) of the algorithm for combined processing based on piecewise

20 constant models.

The case of local intervals without sliding was considered, with the number of points N, $N_{1j} = N(j - 1)$, $N_{2j} = N_{1j} + N - 1$, j = 1, ..., m, $mN = N_f$. The local estimates $H_{n_j}^\circ = c_j^\circ$, j = 1, ..., m were calculated, the sequences of piecewise constant estimates $H_{L_n}^\circ(Ti)$, n = 1,2,3, corresponding to the first processing stage, were formed.

For a sliding local intervals case with the number of points N, the sliding step N_d was selected, as well as the number of sliding local intervals m_0 . Local estimates $H_{n_j}^\circ = c_j^\circ$, $j = 1, ..., m_0$, and the estimates of functions $H_{n_j}^\circ(Ti)$, n = 1,2,3,





 $i = 0, 1, ..., N_f - 1$, corresponding to the first processing stage, were calculated; next, the second stage was performed, where the estimates $H_{S_n}^{\circ}(Ti)$, n = 1,2,3, were found.

For testing, the following values were selected: $N_f = 96$, N = 12, m = 8, $N_d = 1$, $m_0 = 84$, $\sigma = 1.0$ nT, $\sigma_0 = 0.5$ nT. On Fig.1, the calculation results are displayed for H_{MG_1} (Fig. 2a), H_{MG_2} (Fig. 2b), H_{MG_3} (Fig. 2c).; dashed lines with index 1

5 depict the initial functions for the projections of the geomagnetic field intensity vector $H_{MG_n}(Ti)$; lines with index 2 represent the noisy measurements of the projections of the geomagnetic field intensity $H_{M_n}(Ti)$; piecewise constant lines with index 3 represent the results of the first stage $H^{\circ}_{L_n}(Ti)$; solid lines with index 4 show the estimates $H^{\circ}_{S_n}(Ti)$ - the second stage results with weighted averaging.

The performed testing of the processing algorithm on model data for a number of parameters led to the conclusion that the second stage of processing reduces the RMS of the errors of the first stage by 60-80% on average.

4.2 Testing on real observatory geomagnetic data

The developed algorithm was tested using combined processing of 1-second geomagnetic measurements from the INTERMAGNET observatory MBO (Mbour, Senegal). The measurements were recorded on 02.01.2014, they began at 01:16:37 UT, the length of a test fragment was 96 seconds (N_f =96). For the processing algorithm, it was assigned N = 12,

15 and the sliding step $N_d = 1$.

On Fig.2 the test results are shown for H_1 (Fig. 2a), H_2 (Fig. 2b), H_3 (Fig. 2c). Dashed lines with index 1 depict the observatory measurements of the geomagnetic vector projections $H_n(Ti)$; piecewise constant lines with index 2 are related to the first processing stage – the functions for piecewise constant estimates $H_{L_n}^{\circ}(Ti)$ without sliding are displayed; index 3 stands for the line corresponding to the result of the second processing stage – the estimate with weighted averaging $H_{S_n}^{\circ}(Ti)$

20 with sliding.

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The results of testing of the algorithm for combined processing of measurements of projections and absolute magnitudes of geomagnetic field intensity, displayed on Fig.1, 2, proved its satisfactory performance.

5 Error estimation for the algorithm for combined processing of measurements

The estimates of errors of the proposed algorithm for combined processing were found using statistical modelling. The first stage of combined processing was analyzed.

The initial functions for the intensity vector projections were assumed to be constant on a local interval. The values $H_n(k, m, H_0, Ti) = H_n(k, m, H_0)$, i = 0, 1, ..., N - 1, n = 1, 2, 3 were found using the following equations:

 $H_1(k, m, H_0) = H_0 \sin\vartheta_m, H_2(k, m, H_0) = H_0 \cos\vartheta_m \sin\varphi_k, H_3(k, m, H_0) = H_0 \cos\vartheta_m \cos\varphi_k.$

Here H_0 is an assigned absolute magnitude value, φ_k , ϑ_m are the azimuthal and zenithal angles, $\theta_m = \Delta \vartheta m$, $\Delta \vartheta = 2\pi/M_0$, m, k- integer parameters, $m = 0, 1, ..., M_0 - 1$, $\varphi_k = \Delta \varphi k$, $\Delta \varphi = 2\pi/K_0$, $k = 0, 1, ..., K_0 - 1$.





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For all possible values of indices k, m, the realizations of sequences of model normally distributed random numbers with zero mathematical expectation $W_{n,s}(k, m, Ti)$ were formed, where i = 0, 1, ..., N - 1, $n = 0, 1, 2, 3, s = 1, ..., S_0$ is the number of realization for statistical modelling. For n = 0 the variance determining the noise error level for a scalar magnetometer assumed the value σ_0^2 , for n = 1,2,3 the noise error variances for a vector magnetometer assumed the value σ^2 . For $H_n(k, m, H_0)$, $W_{n,s}(k, m, Ti)$ and S_0 random realizations were constructed:

 $H_{n,s}(k, m, H_0, i) = H_n(k, m, H_0) + W_{n,s}(k, m, Ti), i = 0, 1, ..., N - 1, n = 0, 1, 2, 3, s = 1, ..., S_0.$ The results of the algorithm implementation for the first – the estimates $H_{n,s}^{\circ}(k, m, H_0, N, \sigma, \sigma_0)$, $n=1, 2, 3, k=0, 1, ..., K_0-1$, $m = 0, 1, ..., M_0$ -1, $s = 1, ..., S_0$ were calculated, depending on the parameters H_0, N, σ, σ_0 . The error of the processing algorithm $\varepsilon_n^2(k, m, H_0, N, \sigma, \sigma_0)$ was found using averaging over the number of realizations for fixed n, k, m

$$\varepsilon_n^2(k, m, H_0, N, \sigma, \sigma_0) = \frac{1}{s_0} \sum_{s=1}^{s_0} (H_n(k, m, H_0) - H_{n,s}^\circ(k, m, H_0, N, \sigma, \sigma_0))^2.$$
(10)

The error described by Eq. (10) was averaged over the numbers of the absolute geomagnetic vector projections n = 1,2,310 and then over different k, m. The final formula for estimating the error was the following:

$$\varepsilon_f^2(H_0, N, \sigma, \sigma_0) = \frac{1}{\kappa_0 M_0} \sum_{k=1}^{K_0} \sum_{m=1}^{M_0} \frac{1}{3} \sum_{n=1}^{3} \varepsilon_n^2(k, m, H_0, N, \sigma, \sigma_0).$$
(11)

The results of the combined processing algorithm for the first stage were compared with the results of the operation of a possible linear filtering algorithm that was separately applied to the recordings of vector magnetometer channels. The linear filtering algorithm in this case was represented by standard equations:

$$H_{1n,s}^{\circ}(k,m,H_0,N,\sigma) = \frac{1}{N} \sum_{i=0}^{N-1} H_{n,s}(k,m,H_0,i).$$
⁽¹²⁾

The error estimate ε_{1f}^2 (H_0 , N, σ) for the linear filtering algorithm was calculated similar to Eq. (10), (11)

$$\varepsilon_{1n}^{2}(k,m,H_{0},N,\sigma) = \frac{1}{S_{0}} \sum_{s=1}^{S_{0}} (H_{n}(k,m,H_{0}) - H_{1n,s}^{\circ}(k,m,H_{0},N,\sigma))^{2},$$

$$\varepsilon_{1f}^{2}(H_{0}, N, \sigma) = \frac{1}{K_{0}M_{0}} \sum_{k=1}^{K_{0}} \sum_{m=1}^{M_{0}} \frac{1}{3} \sum_{n=1}^{3} \varepsilon_{1n}^{2}(k, m, H_{0}, N, \sigma).$$
(13)

20 The efficiency of the proposed algorithm for combined processing of measurements was estimated using the introduction of a relative decrease factor for the RMS error values $\rho(H_0, N, \sigma, \sigma_0)$

$$\rho(H_0, N, \sigma, \sigma_0) = \frac{\varepsilon_f^2(H_0, N, \sigma, \sigma_0)}{\varepsilon_{1f}^2(H_0, N, \sigma)}.$$
(14)

For statistical modeling , the following values have been assumed: $H_0 = 10000 \text{ nT}$, N = 5, $M_0 = 50$, $K_0 = 50$, $S_0 = 100$, $\sigma_0 = 0.5, 0.3, 0.1, 0.03$ (indices 1, 2, 3, 4), $\sigma = 0.156, 0.312, 0.625, 1.25, 2.50, 5.00, 10.0$. Fig.3 displays the results of $\rho(H_0, N, \sigma, \sigma_0)$ calculation depending on $\log_2(\sigma/\bar{\sigma}), \bar{\sigma} = 0.156$. The estimates of the introduced factor made it possible to 25 get an idea of the effectiveness of the proposed combined processing of measurements.





Analysis of the graphs shows that, for a fixed value σ , the introduced factor ρ decreases when σ_0 decreases, which is physically understandable. It is also seen that this factor tends to limit values with increasing σ . For a fixed H_0 , the increase of *N* leads to the decrease of the factor ρ . The performed statistical modeling for a wide range of parameters shows that the estimates for ρ are about 0.15-0.3 which indicates the efficiency of the proposed combined processing.

5 6 Conclusions

The proposed method for combined processing of the measurements of projections and absolute magnitudes of geomagnetic field intensity vectors and the corresponding two-stage algorithm developed appeared to be satisfactorily workable. The testing of the developed combined processing algorithm on model and observatory measurement data proved its efficiency. The approach realized in this paper, based on the proposed method, provides the increase of accuracy of the measurements of

10 the projections of the geomagnetic intensity vector.

Statistical modelling for the developed algorithm for combined processing of measurements shows that, at the first stage, the relative decrease factor of RMS errors can reach values of 0.15-0.3 and at the second stage the decrease of the first-stage RMS errors can reach approximately 60-80%.

Further reduction of the RMS noise errors can be implemented based on combined processing using local piecewise linear models for the values of projections and absolute magnitudes of geomagnetic field intensity vectors. The proposed combined processing algorithm can be implemented for many practically important tasks, in particular, when optimizing the operation of three-component accelerometer systems or three-component angular velocity sensors.

7 Author contribution

V.G. Getmanov is the author and the main developer of the method presented in this research. Also the basic work on
 programming was done by him. A.D. Gvishiani performed the optimization and upgrade of the algorithm. R.V. Sidorov collected, analyzed and prepared the geomagnetic data for the computational tests.

8 Competing interests

The authors declare that they have no conflict of interest.

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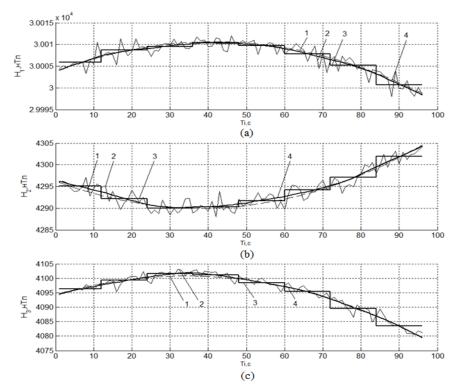


Figure 1: Results of testing of the processing algorithm on model measurement data.

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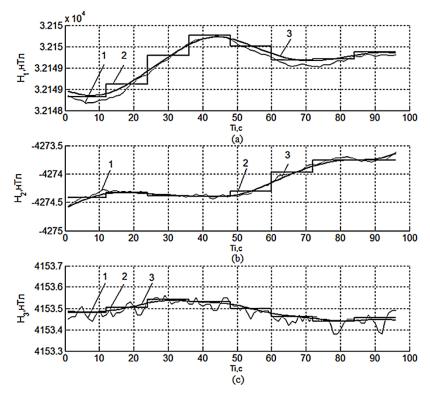


Figure 2: Results of testing of the processing algorithm on observatory measurement data.





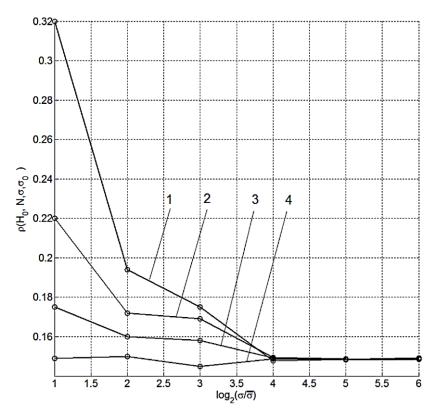


Figure 3: Results of calculation of the relative decrease factors for the errors